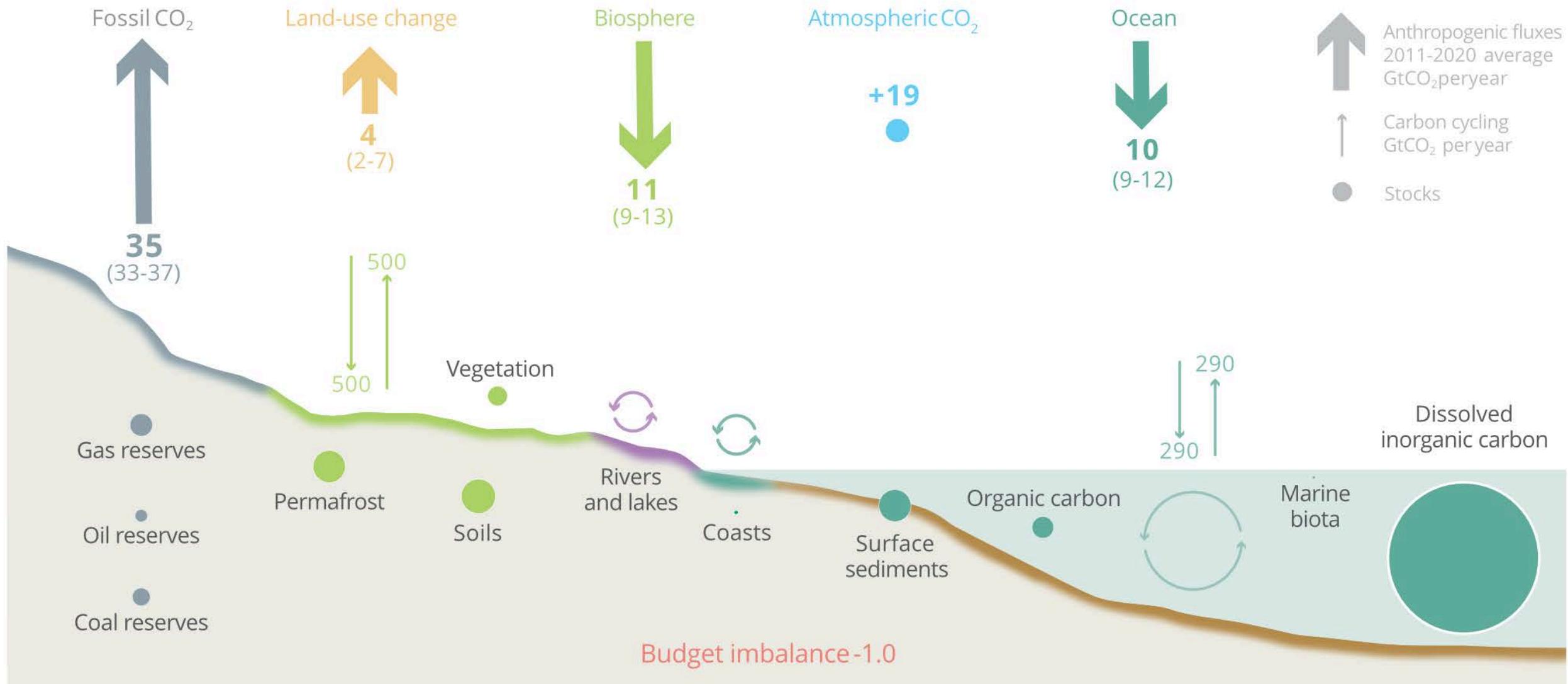


# The importance of SST and in situ observations for improving global air/sea CO<sub>2</sub> flux estimates

T.G. Bell, M. Yang,  
L. Marie, Y. Dong,  
T. Smyth, V. Kitidis,  
I. Brown, J. Shutler,  
G. Tilstone

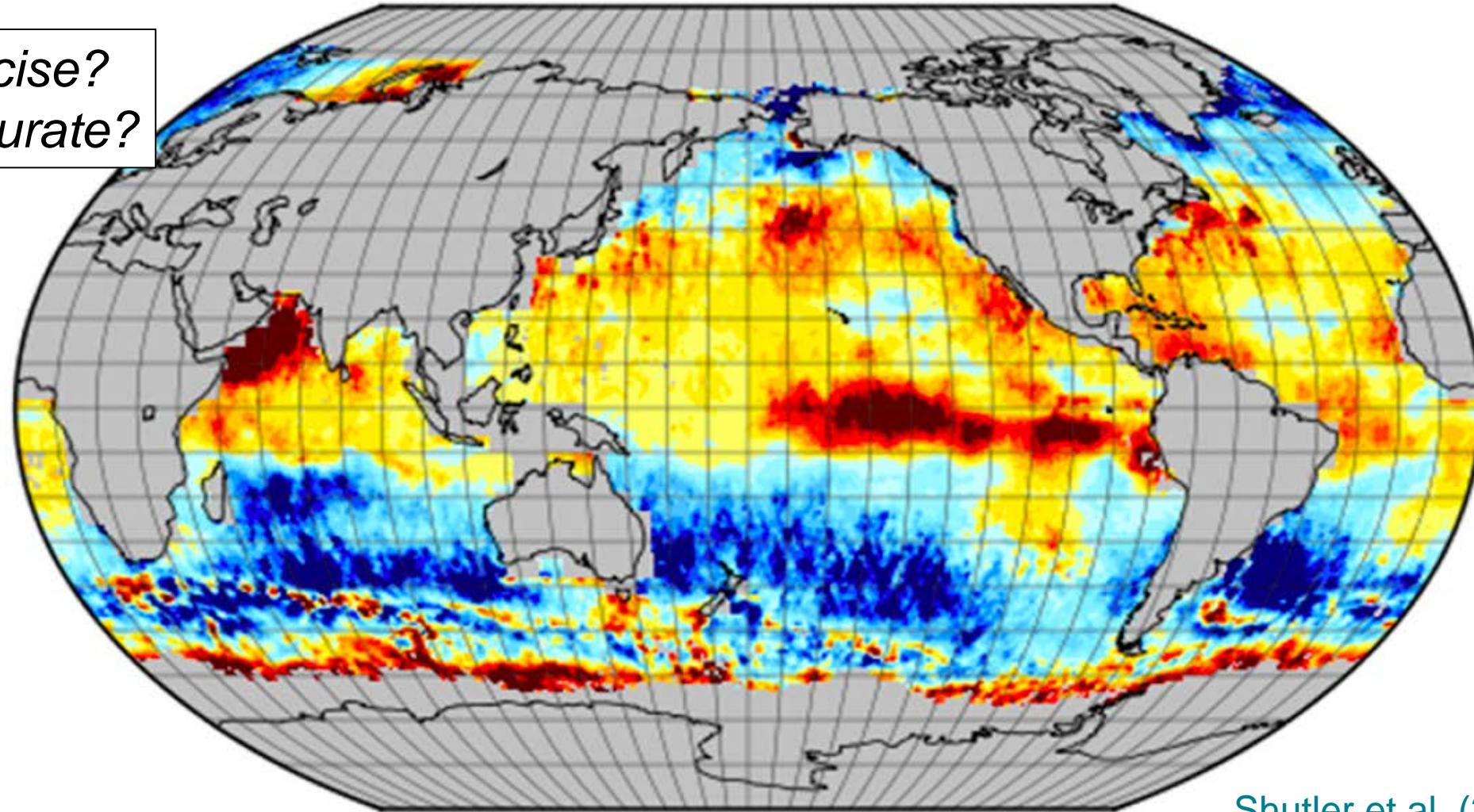


# Why measure air-sea CO<sub>2</sub> fluxes?



# Satellite-based estimates of air-sea CO<sub>2</sub> flux

How precise?  
How accurate?



Shutler et al. (2016)



Air-sea CO<sub>2</sub> flux (gC m<sup>-2</sup> day<sup>-1</sup>) for August 2000

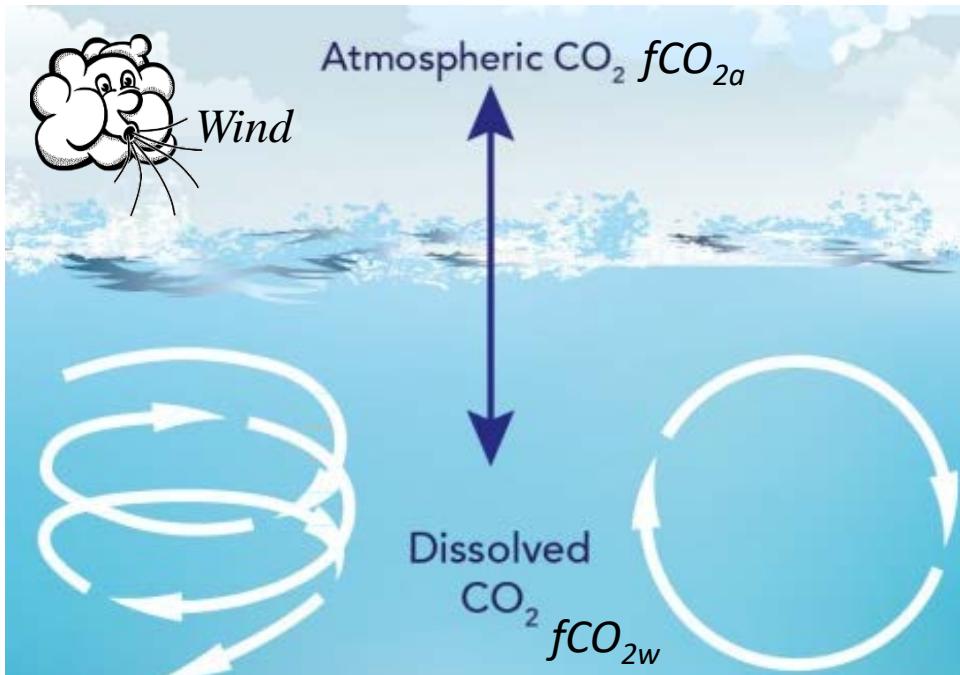
**Indirect measurements** of oceanic/atmospheric variables to calculate flux

$$\text{Flux} = K \cdot \Delta C$$

$$\text{Flux} = K(\alpha_w fCO_{2w} - \alpha_i fCO_{2a})$$

$\alpha$  (solubility), function of temperature

$K$  (gas transfer velocity), function of physical forcing(s)



# Estimating air-sea CO<sub>2</sub> fluxes

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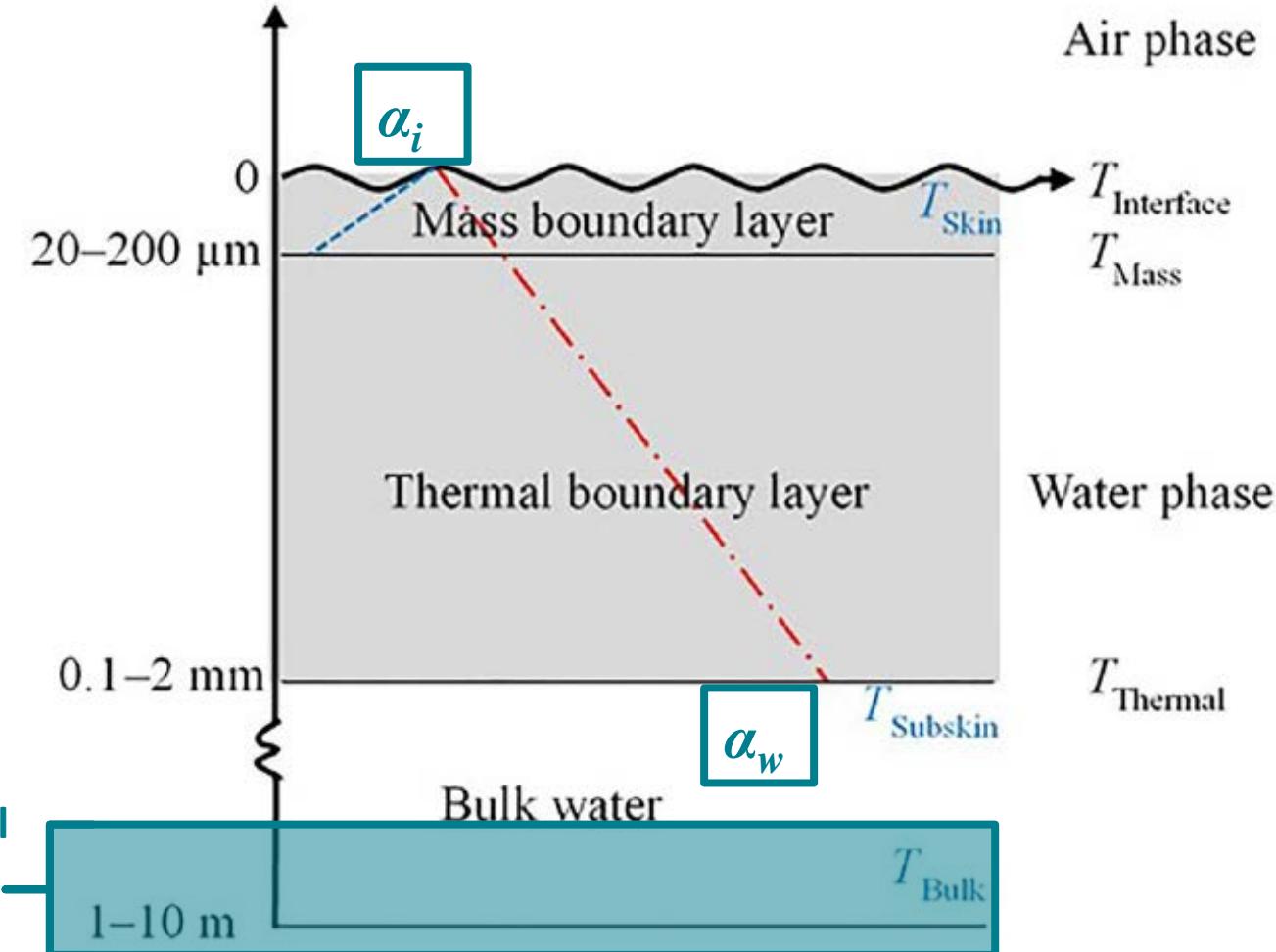
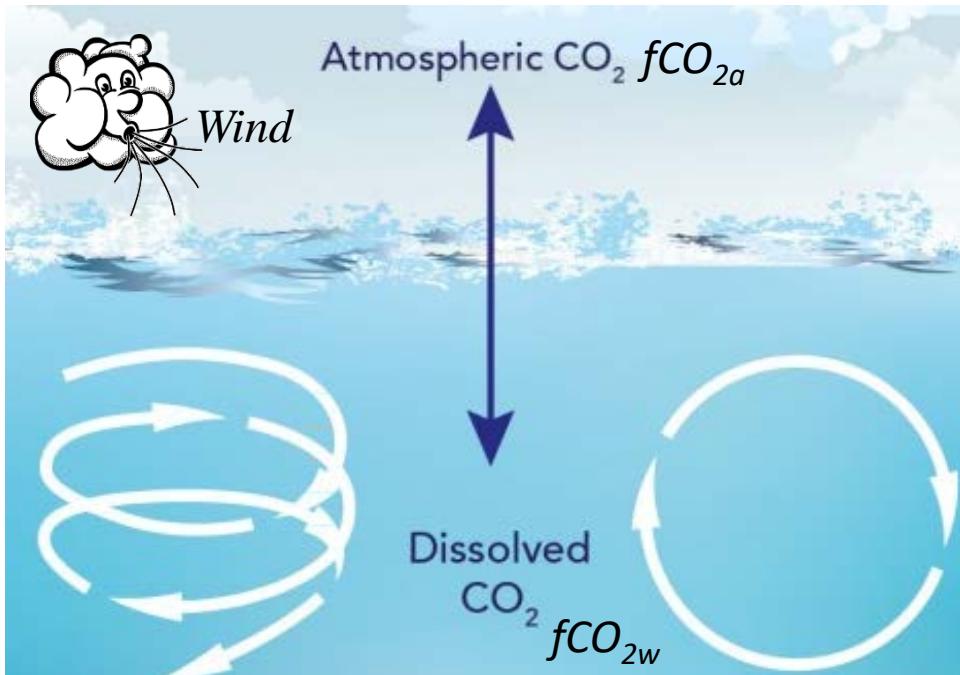
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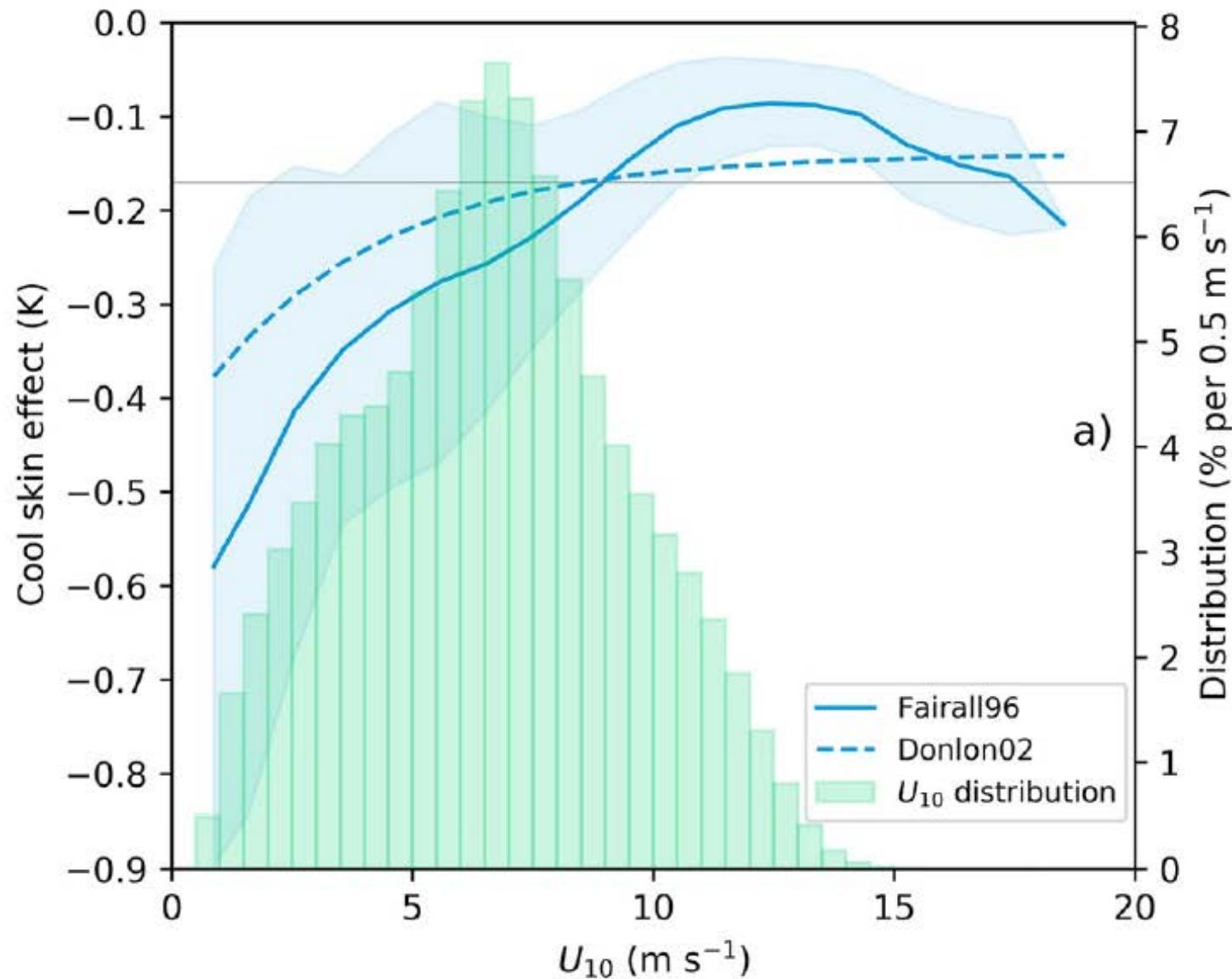
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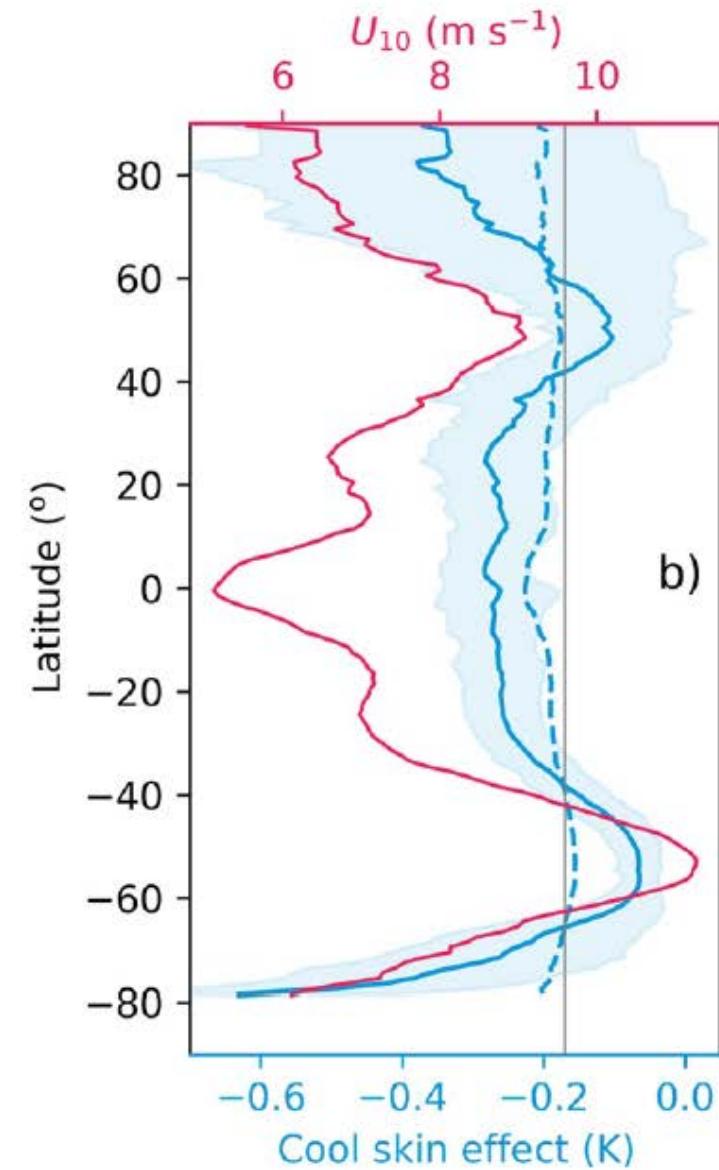
....AND seawater warms inside ships!

$$fCO_{2w} = fCO_{2w, \text{equ}} e^{0.0423(T_{w,bulk} - T_{\text{equ}})}$$

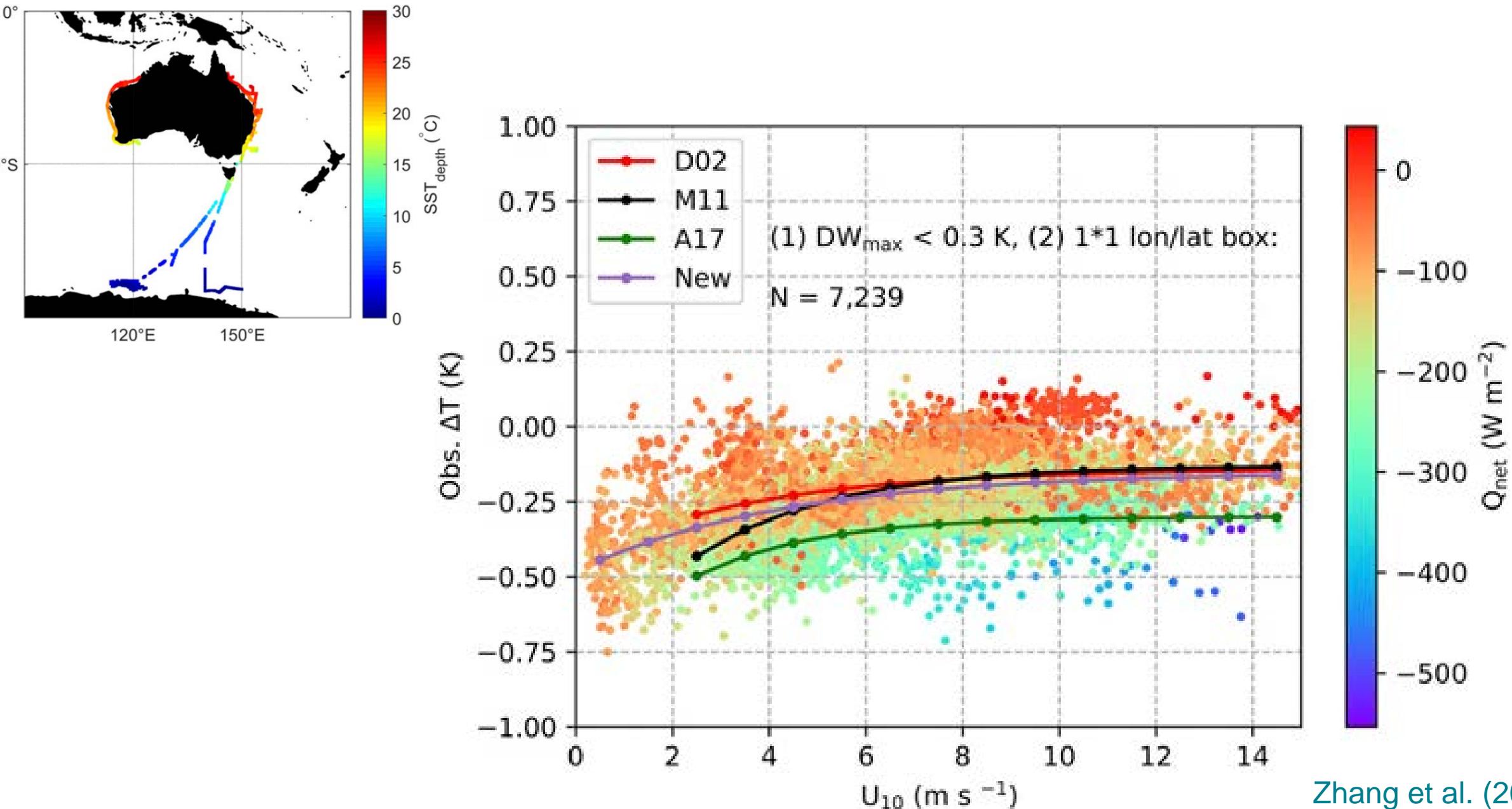
# Choice of cool skin model



Dong et al., GBC (2022)



# ISAR cool skin observations



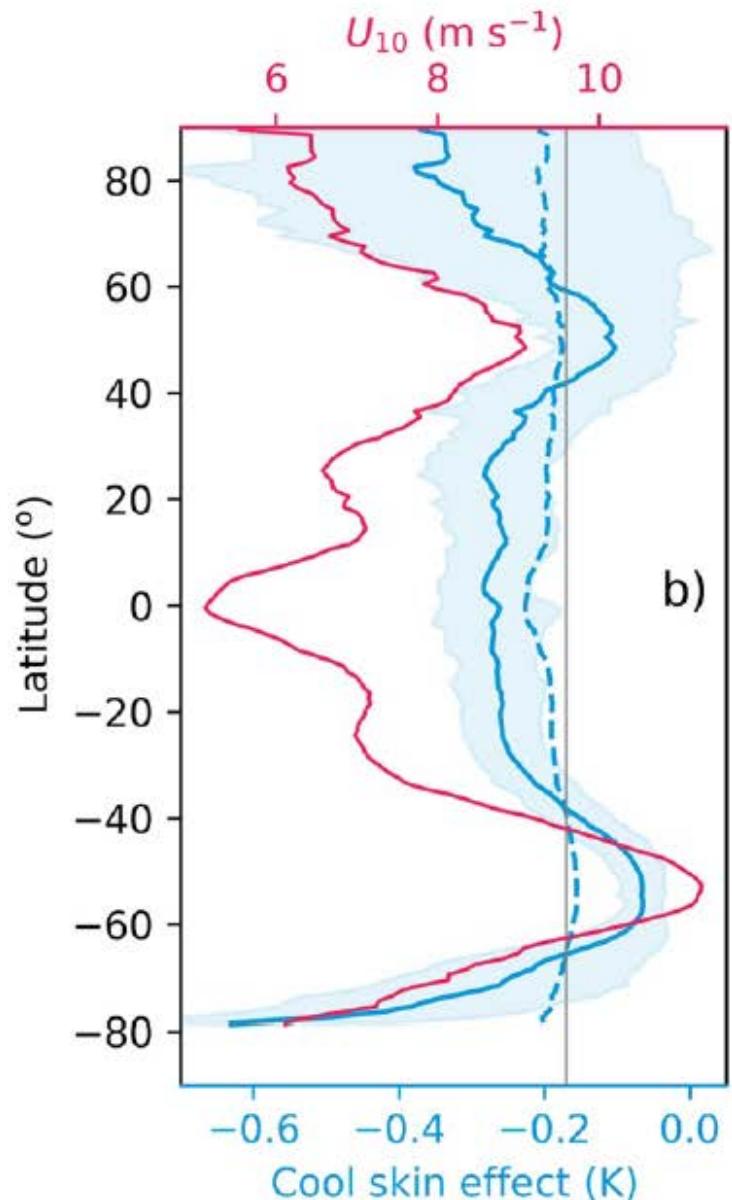
# Is the cool skin correction important?

Dong et al., GBC (2022):

- Global ocean uptake estimate:

No temperature correction	$1.7 \pm 0.4 \text{ Pg C yr}^{-1}$
With temperature corrections	$2.2 \pm 0.4 \text{ Pg C yr}^{-1}$

- 30% enhancement in carbon uptake!
- Cool skin correction responsible for  $\frac{3}{4}$  of this adjustment
- Independent ocean carbon inventory estimate =  $2.1 \pm 0.4 \text{ Pg C yr}^{-1}$



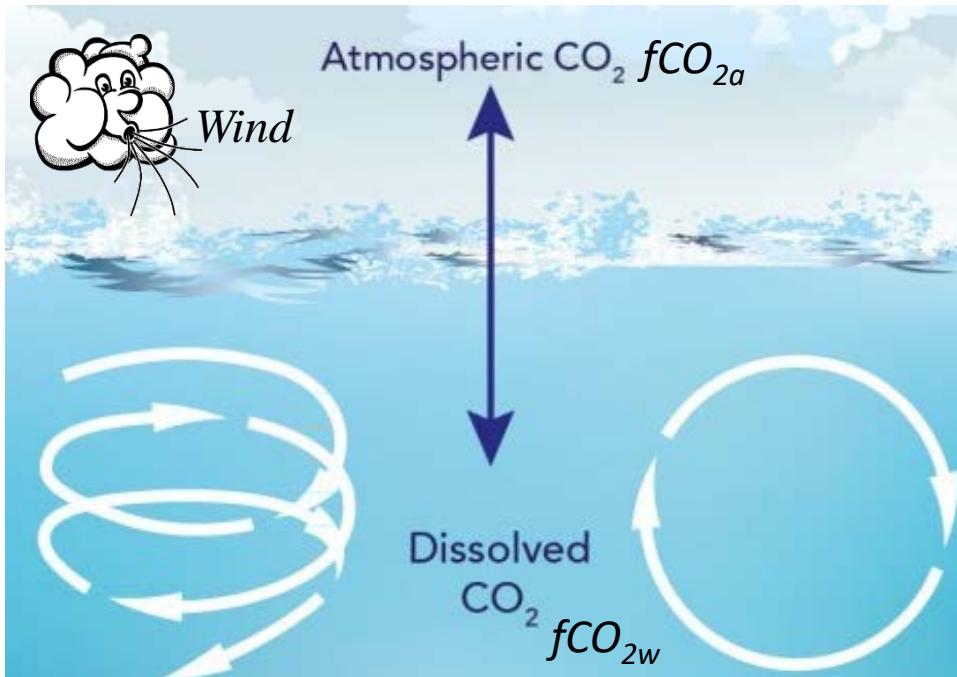
# Estimating air-sea CO<sub>2</sub> fluxes

Indirect measurements of oceanic/atmospheric variables to calculate flux

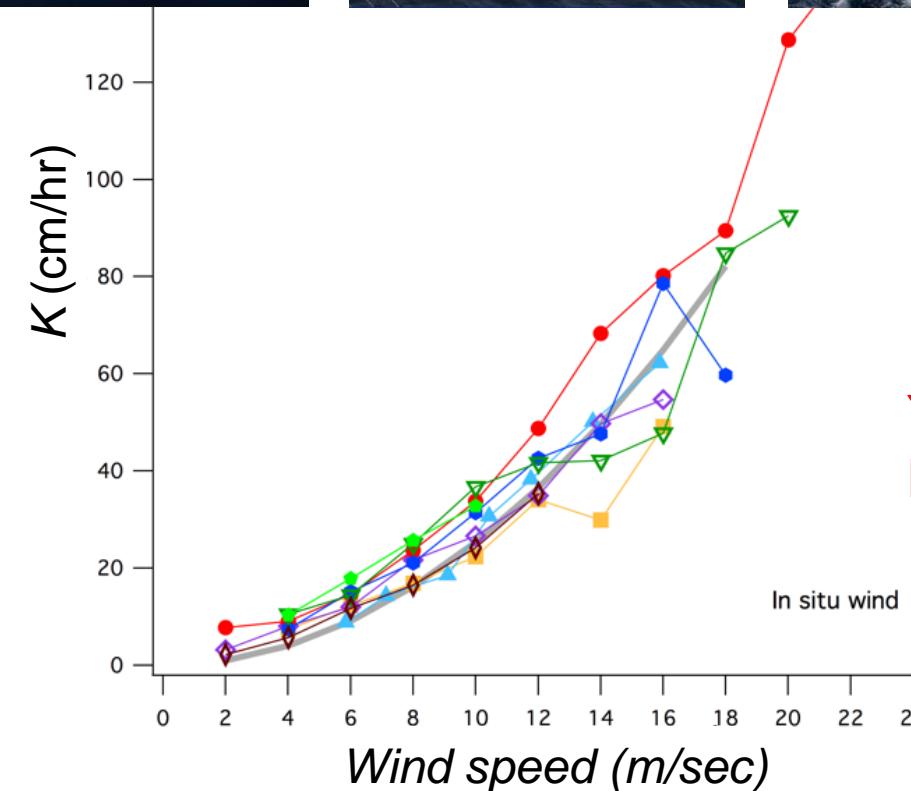
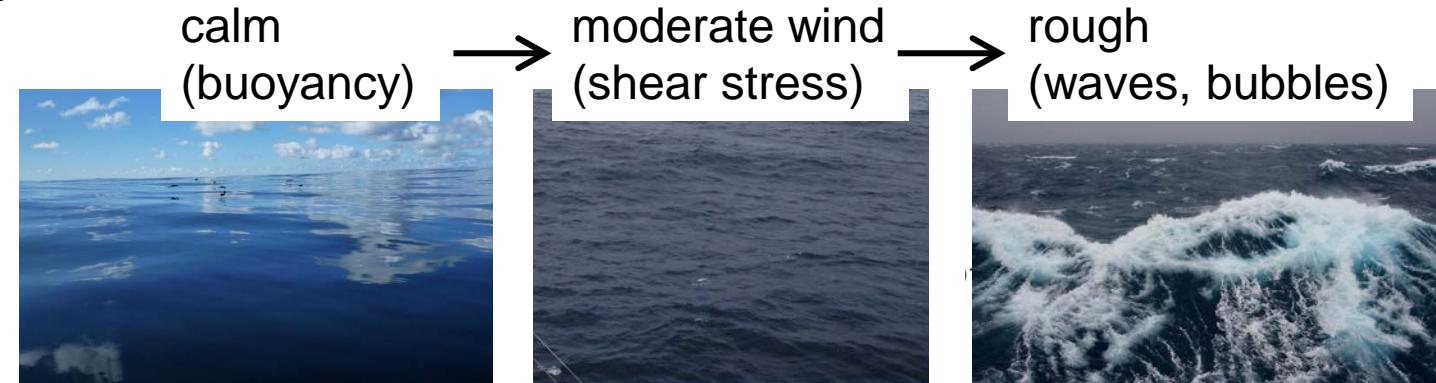
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$\alpha$  (solubility), function of temperature  
K (gas transfer velocity), function of physical forcing(s)



*K* (gas transfer velocity) uncertainty



Yang et al.,  
Frontiers (2022)

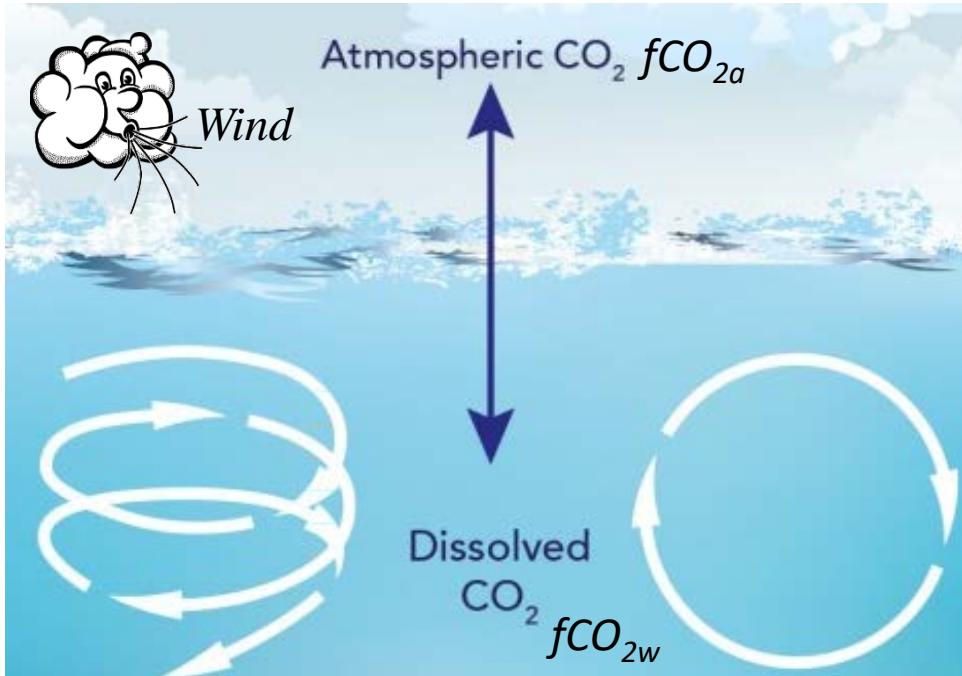
# Estimating air-sea CO<sub>2</sub> fluxes

Indirect measurements of oceanic/atmospheric variables to calculate flux

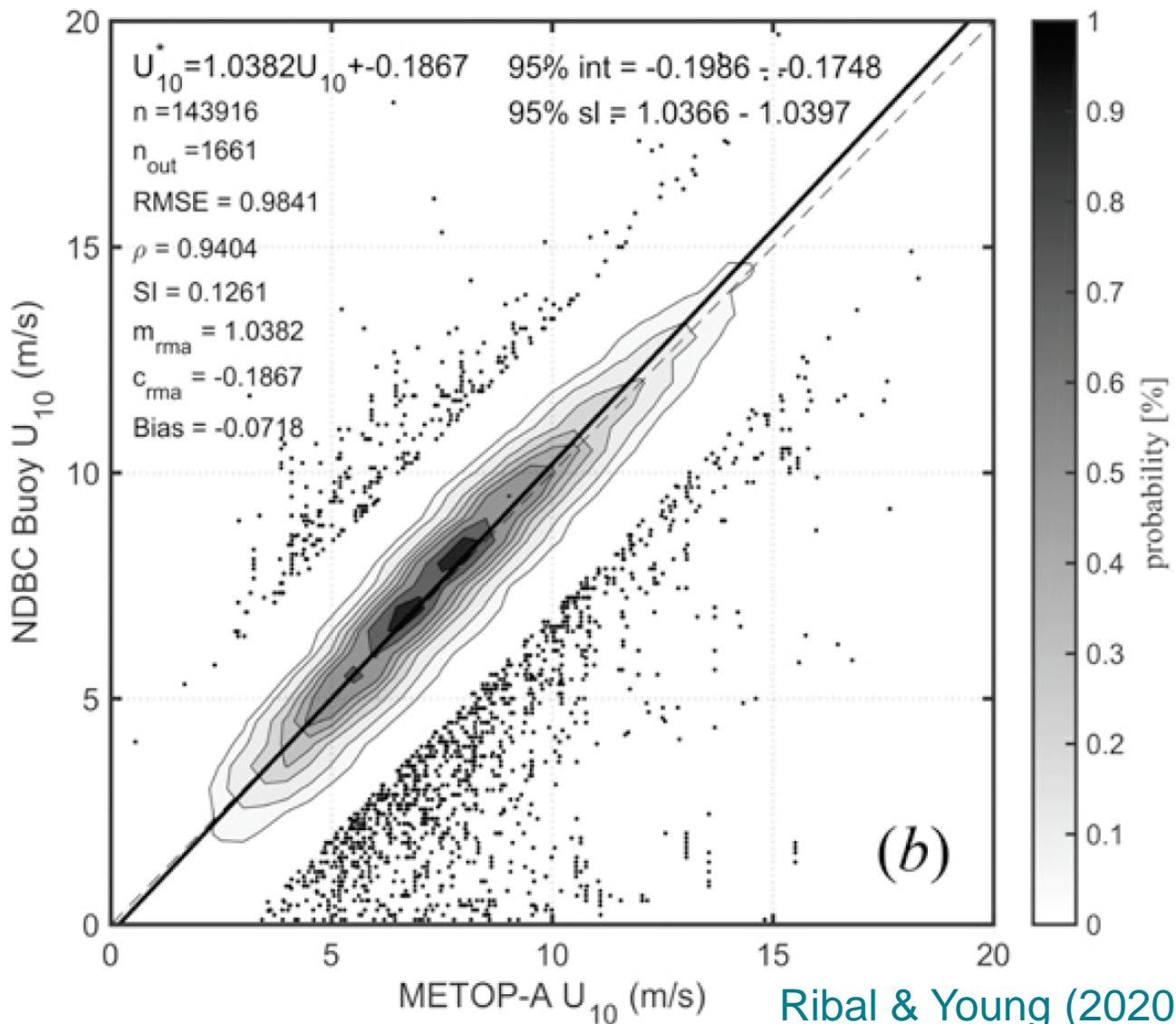
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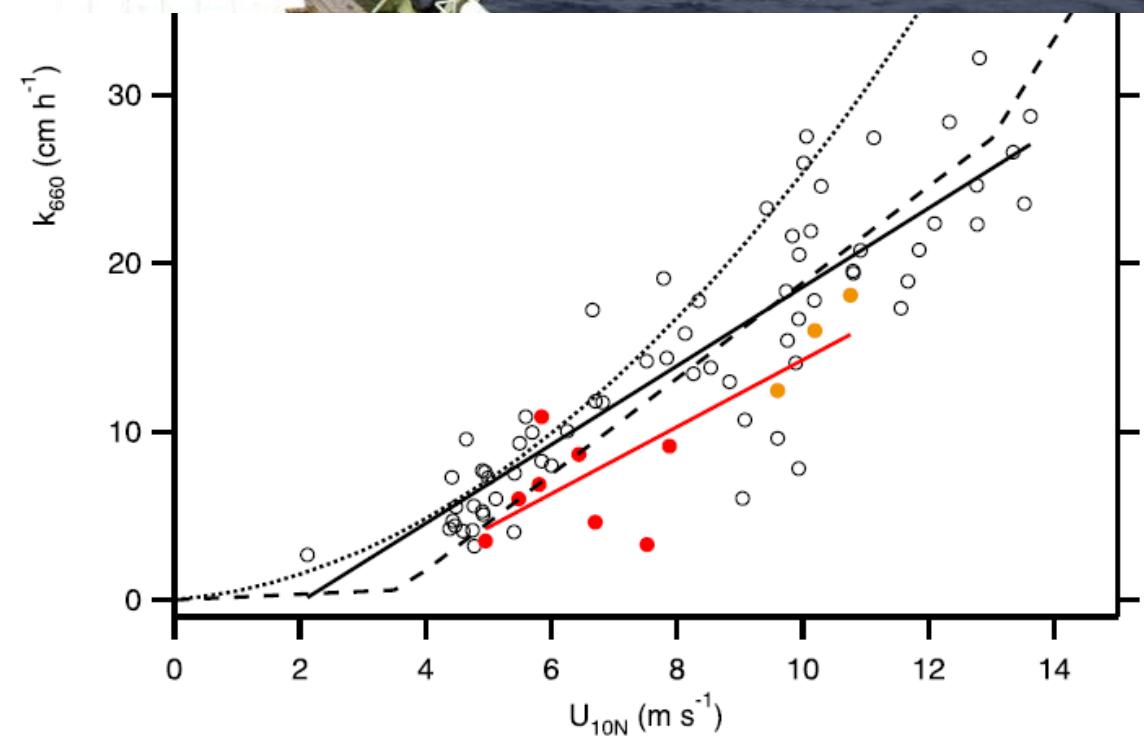
$\alpha$  (solubility), function of temperature  
K (gas transfer velocity), function of physical forcing(s)



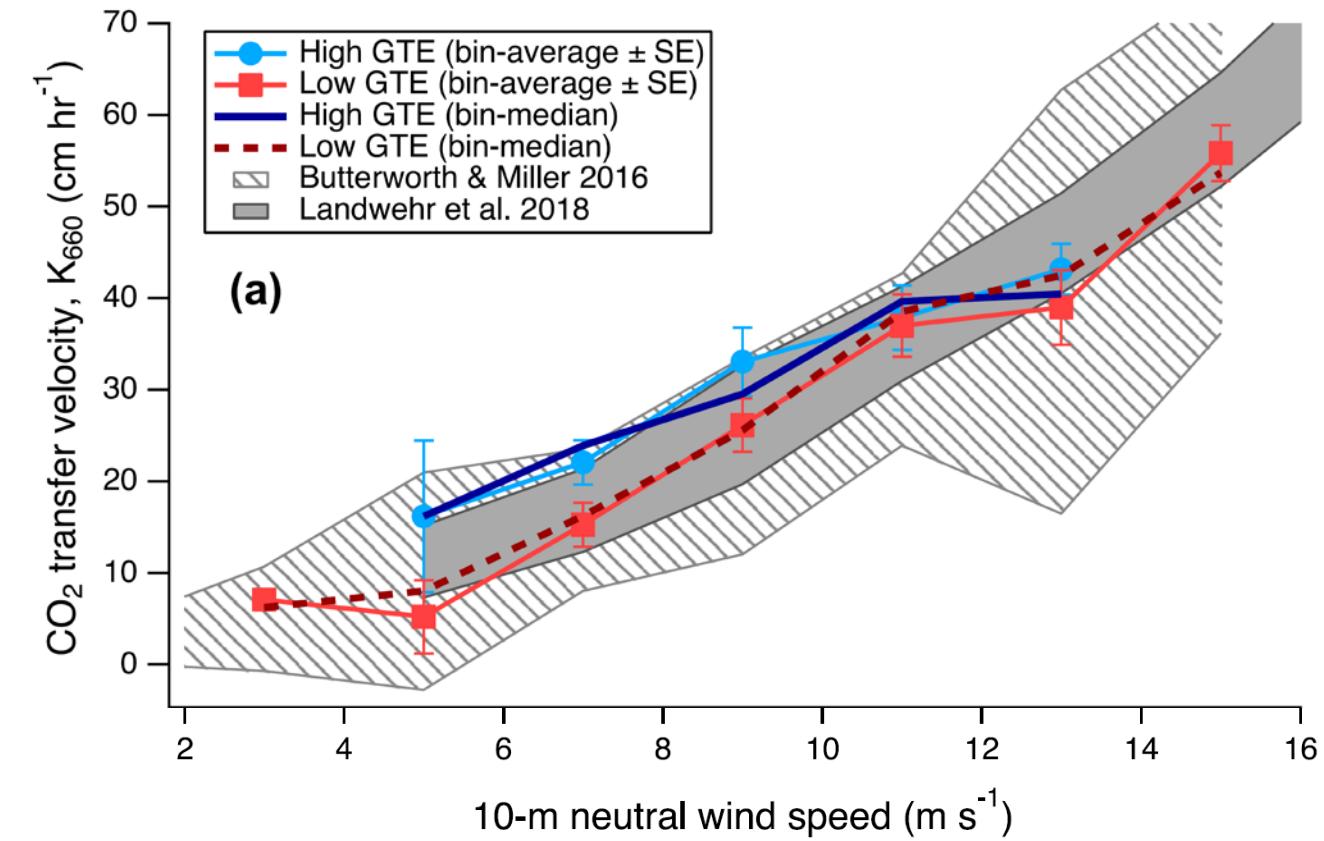
Satellite wind speed ( $U_{10}$ ) uncertainty



# Impact of surfactants on sea surface



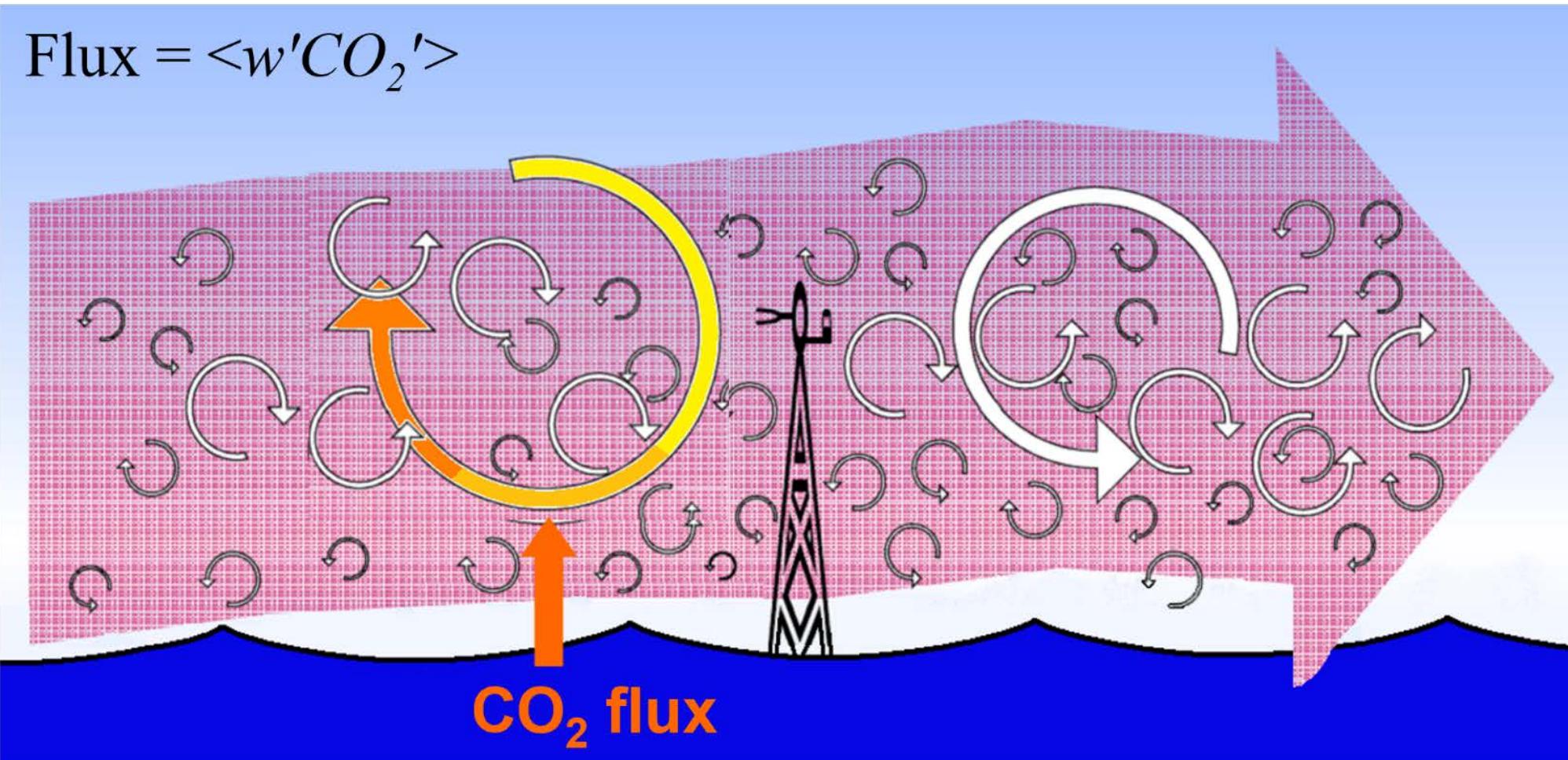
Salter et al. (2011)



Yang et al., Nature Sci. Rep. (2021)

# Measuring air-sea CO<sub>2</sub> fluxes

**Direct measurements** (eddy covariance) – independent estimate of flux



Combined observations are a tool to investigate processes:  $K = \text{Flux} / (\alpha_w fCO_{2w} - \alpha_i fCO_{2a})$

# Autonomous air-sea CO<sub>2</sub> flux systems



Flux observations:

- Closed path
- Dried
- Licor/Picarro

See Dong et al., ACP (2021)

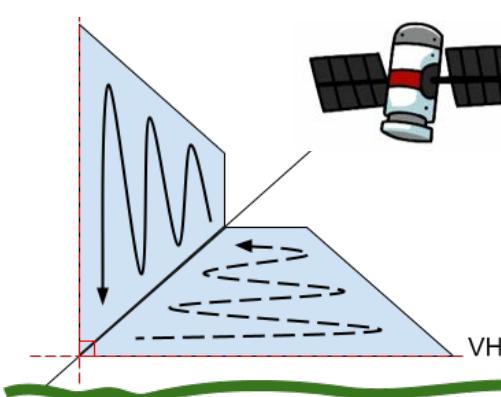
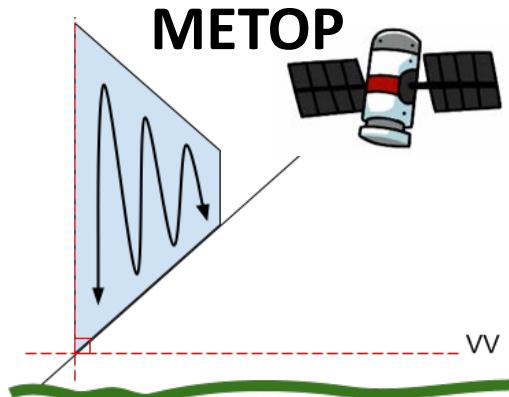


Seawater pCO<sub>2</sub>

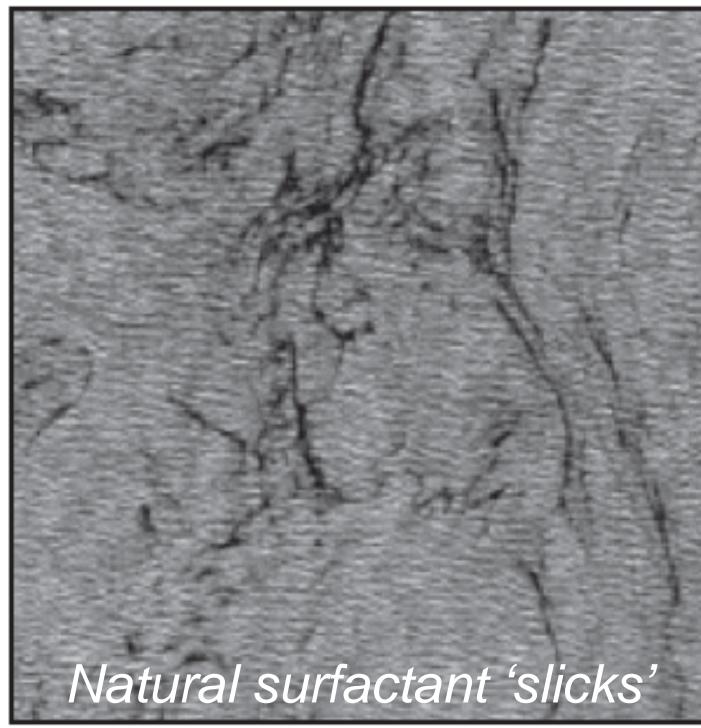
Eddy covariance CO<sub>2</sub> fluxes



# Sea surface scattering



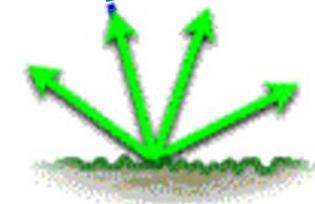
METOP  
Sentinel 1



Wang et al. (2019)



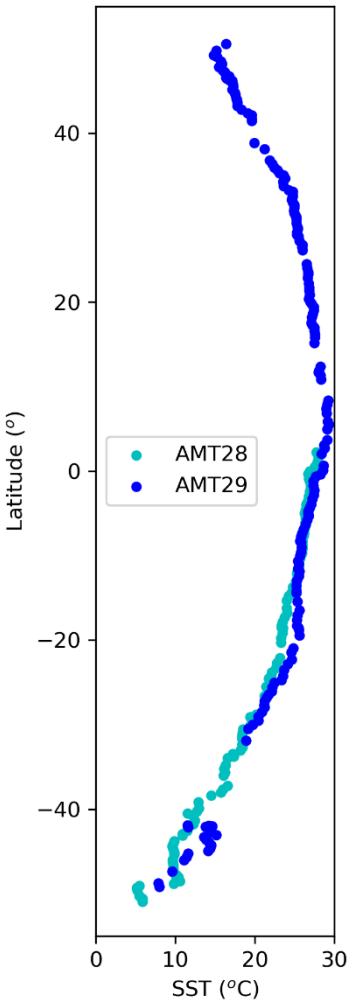
Kudryavtsev  
et al. (2014)



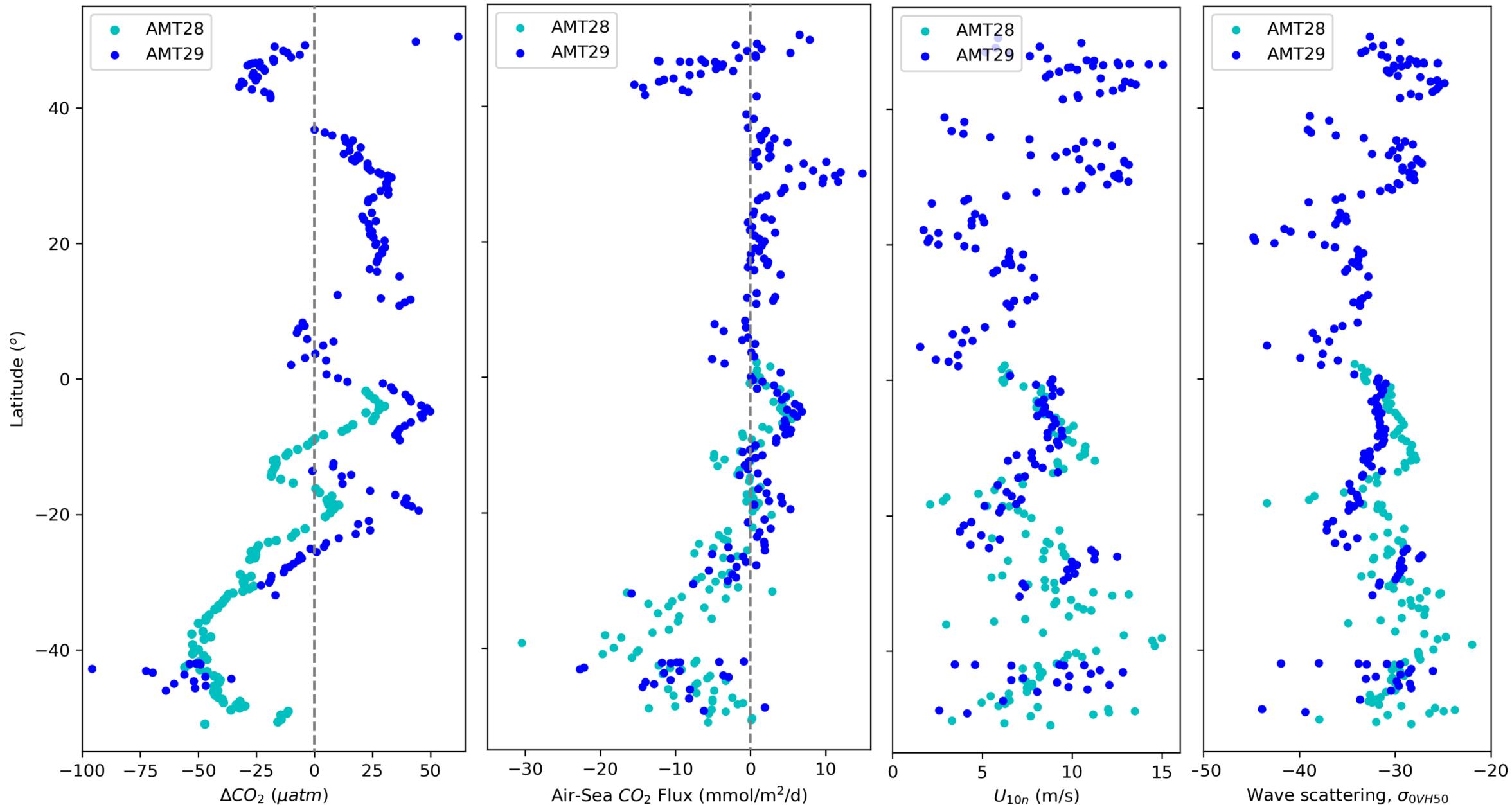
- Scattering,  $\sigma_0$  (C-band radar)
- Cross Polarisation (Horizontal/Vertical)
- Angles:  $30^\circ$ - $50^\circ$
- Measures roughness (waves, whitecaps, etc.)

# Atlantic Meridional Transect (AMT)

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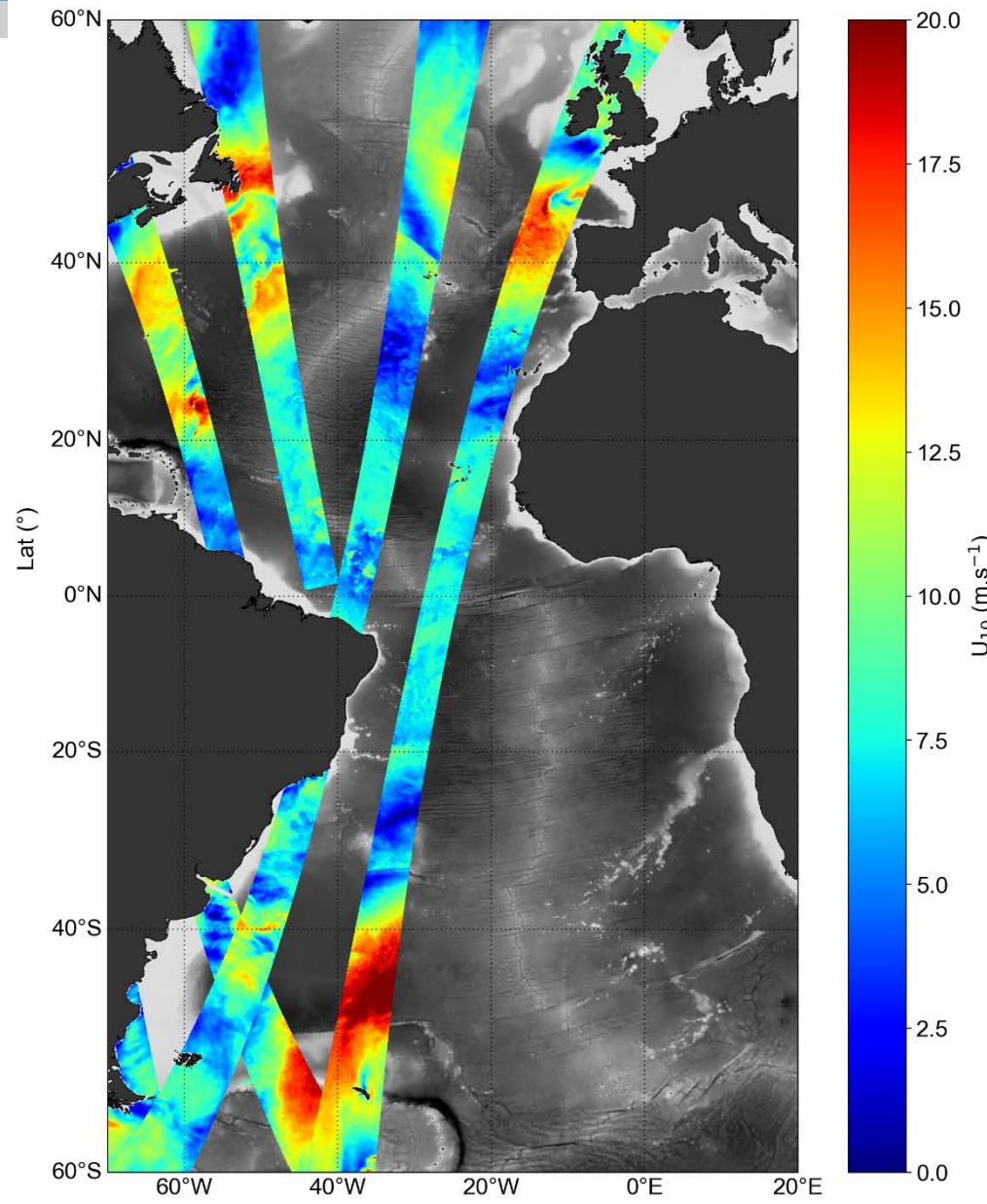
## Latitudinal trends



- Cool skin temperature correction enhances global CO<sub>2</sub> uptake estimates by >20%
- Theoretical adjustment though! Needs verification
- K vs *in situ* radar backscatter ( $\sigma_0$ ) relationship similar strength to K vs wind speed relationship
- Cross polarization  $\sigma_0$  (VH) and 50° angle are optimal
- Reasonable comparison between satellite  $\sigma_0$  and *in situ* observations
- In situ K vs. METOP  $\sigma_0$  VV
  - Few observations
  - Corroborates *in situ*  $\sigma_0$  and K relationship

**Future:** - More *in situ* observations needed

- Investigation of satellite  $\sigma_0$  potential



METOP: 22<sup>nd</sup> Nov. 2019